

Abstract

- Ice, Cloud and land Elevation Satellite-2 (ICESat-2) sea ice height (SIH) data are used to calculate 10 m neutral drag coefficients using parameterizations by Garbrecht et al. 2002.
- **Motivation:** assessing the impact of sea ice topography on the air-ice energy/momentum exchange has thus far been limited to airborne measurements that are both temporally and spatially limited; here we try to extend these methods via remote sensing for large-scale analysis of the effect of sea ice roughness.

Hypothesis

ICESat-2 sea ice height data can be used to extrapolate drag coefficients derived from airborne topography data for an accurate pan-Arctic study of atmosphere-ice momentum transfer

1. Formulation

Adding together skin drag $C_{dn10,i}$ (small-scale roughness) and form drag $C_{dn10,f}$ (distinct obstacles e.g. ridges), the total 10 m neutral drag coefficient C_{dn10} may be calculated as

$$C_{dn10} = C_{dn10,i} + C_{d10,f}$$

- skin drag $C_{dn10,i}$ is kept constant ($8.38 \cdot 10^{-4}$) and form drag is a function of obstacle height H_e and how far apart the obstacles are spaced x_e .
- c_w is the coefficient of resistance estimated as $0.185 + 0.147H_e$.
- z_0 is the surface roughness length over smooth ice given below.

$$C_{d10,f} = \frac{c_w H_e}{\pi x_e} \frac{\left[\ln \left(\frac{H_e}{z_0} \right) - 1 \right]^2 + 1 - 2 \frac{z_0}{H_e}}{\left[\ln \left(\frac{10}{z_0} \right) \right]^2}$$

$$z_0 = 10 \exp \left(- \frac{\kappa}{\sqrt{C_{dn10,i}}} \right)$$

$\kappa = v$. Karman constant (0.4)

Formulas retrieved from Castellani et al. 2014

2. ICESat-2 Data

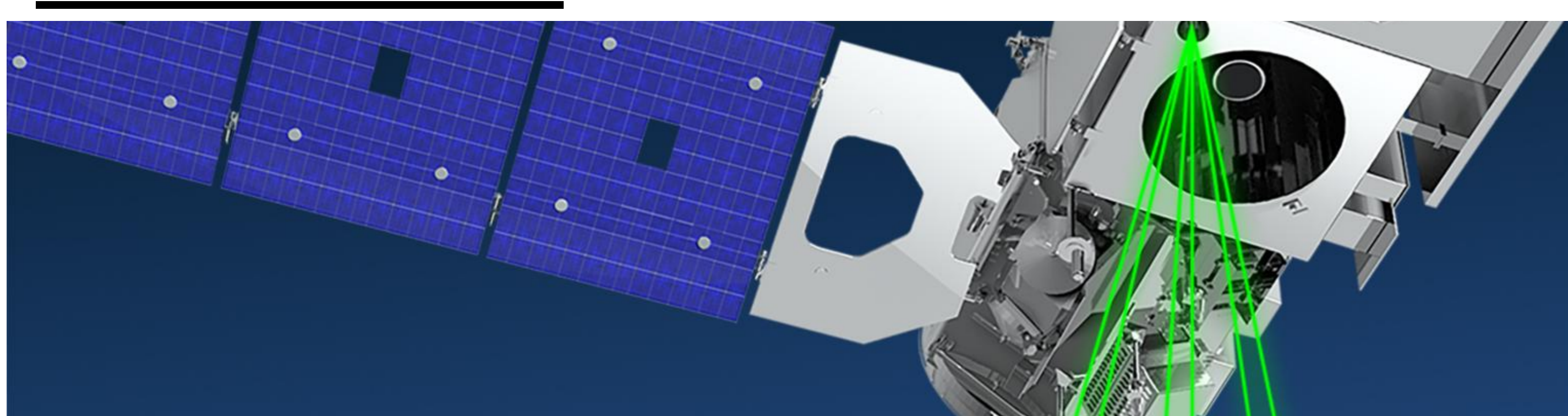


Fig. 1: Illustration of NASA's Ice, Cloud and land Elevation Satellite-2 (ICESat-2), a mission to measure the changing height of Earth's ice. Credits: NASA

- The Advanced Topographic Laser Altimeter System (ATLAS) onboard ICESat-2 sends 10,000 laser light pulses every second: transmits laser pulses every 70 cm along the ground track.
- ATL07 (ATLAS data product no. 7) mean sea ice heights are calculated (via a double Gaussian fit) for segments over which 150 signal photons are accumulated.
 - As a result, mean along-track resolution is lowered to 30 m and average point spacing becomes roughly 15 meters.

REFERENCES

- Garbrecht, T., Lüpkes, C., Hartmann, J. and Wolff, M. M. (2002): Atmospheric drag coefficients over sea ice - validation of a parameterisation concept, Tellus 54A(2), pp. 205-219
- Castellani, G., Lüpkes, C., Hendricks, S., and Gerdes, R. (2014), Variability of Arctic sea-ice topography and its impact on the atmospheric surface drag, J. Geophys. Res. Oceans, 119, 6743–6762, doi:10.1002/2013JC009712.

3. Pan-Arctic Drag Coefficient Map

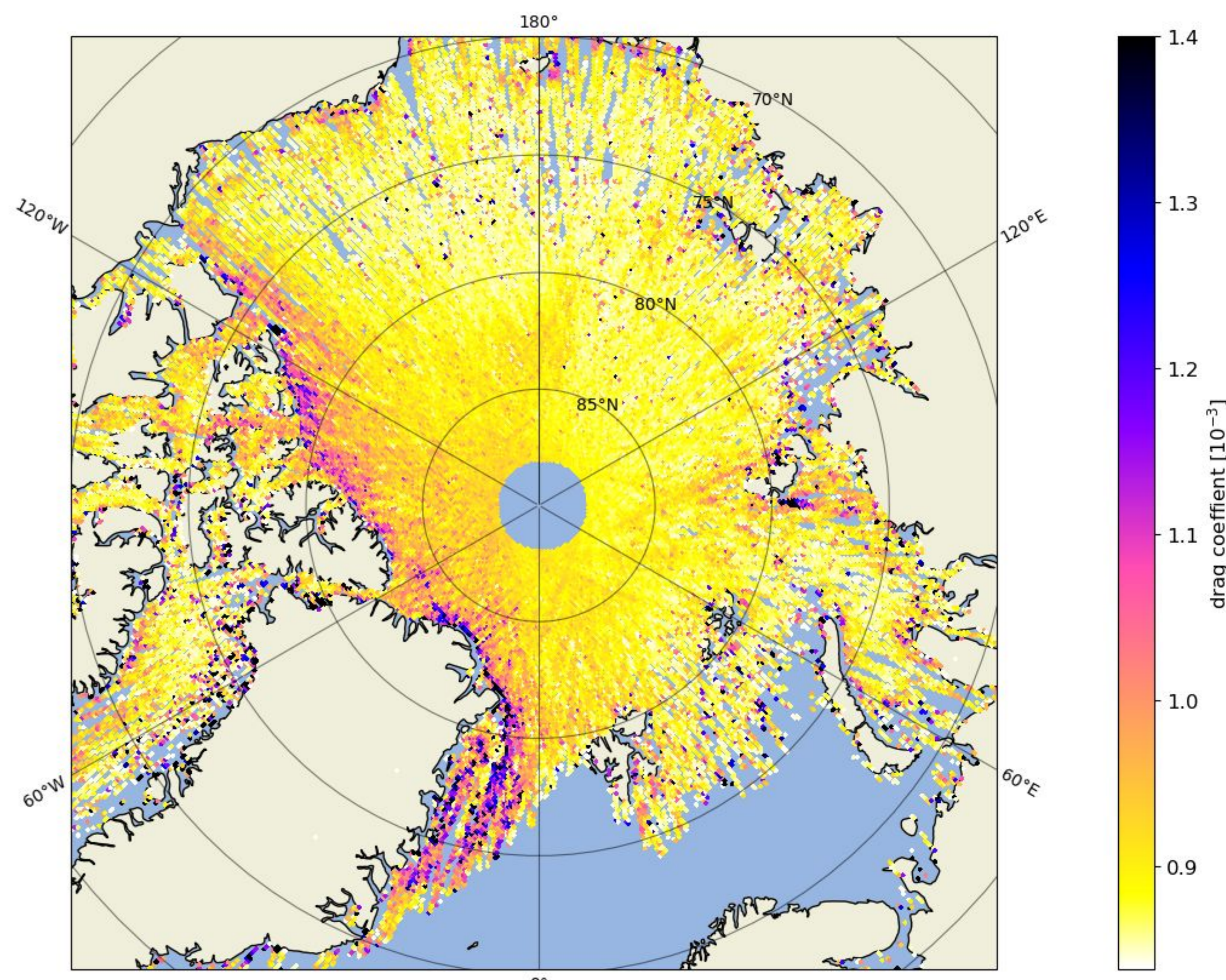


Fig. 2: Pan-Arctic drag coefficients calculated from all ATL07 SIH tracks from March 2020.

4. Comparison with Operation IceBridge

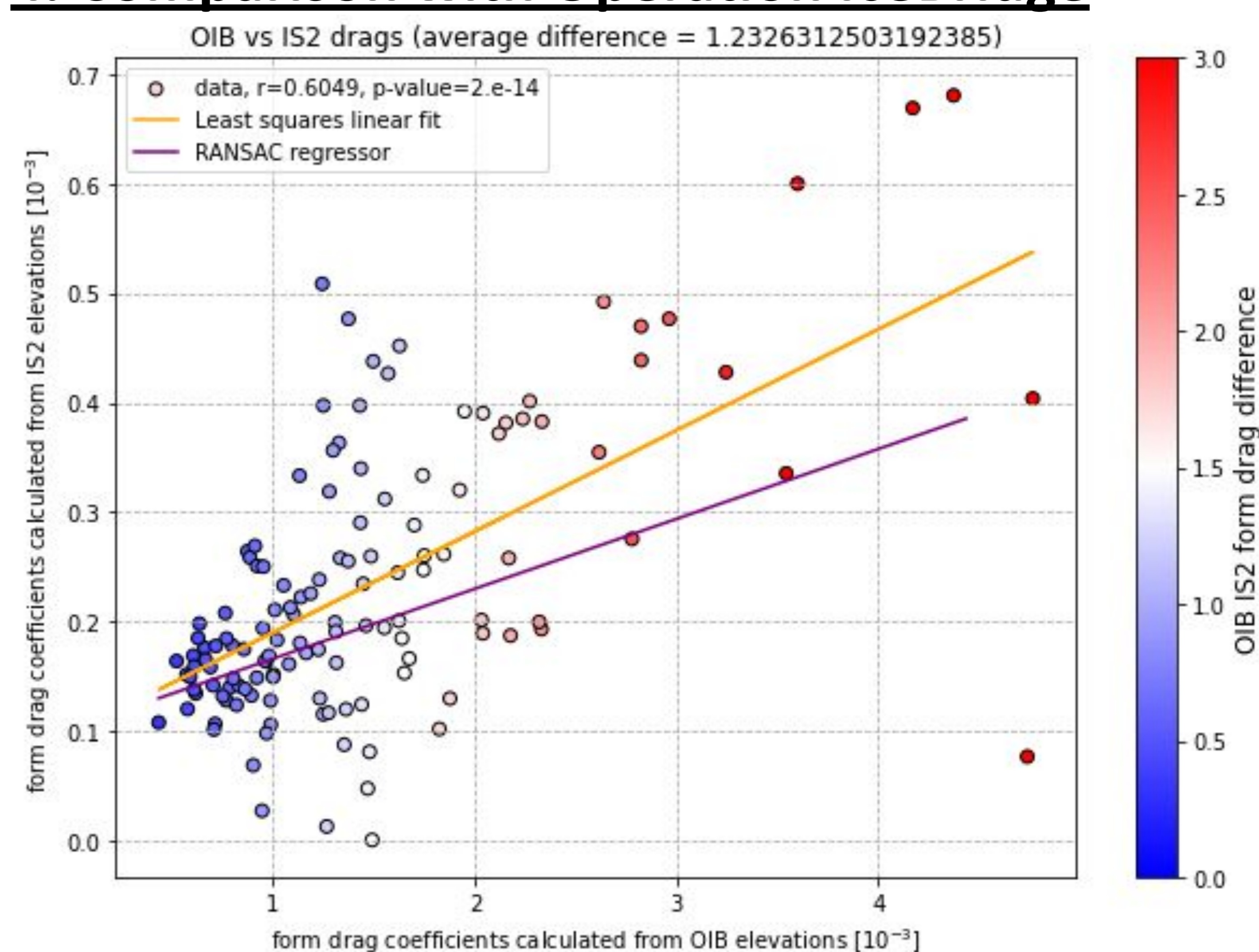


Fig. 3: OIB data-derived and ICESat-2 ATL07-derived form drag coefficients plotted against each other for the time period of April 2019; the colorbar shows the average difference between the two.

5. Conclusion & Outlook

- Using high-precision satellite altimetry, it is possible to retrieve monthly pan-arctic drag coefficients.
 - The retrieved drag coefficients heavily underestimate actual drag due to length of ICESat-2 segments.
- Using additional airborne topography data it should be possible to remedy this underestimation and fine-tune existing parametrization to retrieve more accurate drag coefficients with ICESat-2 SIH data.
- Post-refinement the dataset is meant to be implemented in the coupled Arctic atmosphere-ocean-sea ice model HIRHAM-NAOSIM; where it is expected to improve model simulations.